

## Water relations of adult Norway spruce (*Picea abies* (L) Karst) under soil drought in the Vosges mountains: water potential, stomatal conductance and transpiration

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**Summary** — The effects of soil water depletion on sap flow, twig water potential, stomatal and canopy conductance were analysed in 2 plots of a 30-year-old stand of Norway spruce. One was subjected to an imposed drought; the other was watered by irrigation. Predawn water potential in trees from the dry plot decreased to  $-1.2$  MPa. In the watered plot, a low between-tree variability of sap flux density was observed, with maximum values of  $1.2$ – $1.9$   $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ , corresponding to about  $0.5$   $\text{mm} \cdot \text{h}^{-1}$ . In the dry plot, sap flux density showed a higher variability, and decreased during the summer to a minimum midday value of  $0.05$   $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ . Tree transpiration and stomatal conductance showed a strong reduction in association with drought development, during which the predawn water potential decreased from  $-0.4$  to  $-0.6$  MPa. Canopy conductance was calculated from the reverse of the Penman–Monteith equation assuming that vapour flux over the stand was equal to the estimated stand sap flow. Effects of climatic factors and drought on canopy conductance variations were taken into account in a multi-variable transpiration model.

**transpiration / stomatal conductance / canopy conductance / water potential / drought / sap flow / *Picea abies***

\* Correspondence and reprints.

**Abbreviations:**  $\Psi_t$ : twig water potential (MPa);  $\Psi_{pd}$ ,  $\Psi_m$ : predawn and diurnal minimal twig water potential (MPa), respectively;  $F_d$ : xylem sap flux density ( $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ );  $F$ : total xylem sap flow ( $\text{dm}^3 \cdot \text{h}^{-1}$ ); SA: sapwood area ( $\text{dm}^2$ );  $T_w$ ,  $T_d$ : transpiration of watered and dry plot ( $\text{mm} \cdot \text{h}^{-1}$ ,  $\text{mm} \cdot \text{d}^{-1}$ ), respectively; TM: maximal plot transpiration ( $\text{mm} \cdot \text{h}^{-1}$ ,  $\text{mm} \cdot \text{d}^{-1}$ );  $g_s$ : stomatal conductance ( $\text{cm} \cdot \text{s}^{-1}$ );  $g_c$ : canopy conductance to water vapour ( $\text{cm} \cdot \text{s}^{-1}$ ); VPD: vapour pressure deficit (Pa, hPa);  $R_g$ : global radiation ( $\text{W} \cdot \text{m}^{-2}$ ).

**Résumé — Relations hydriques chez l'épicéa commun (*Picea abies* (L) Karst) soumis à une sécheresse édaphique dans les Vosges : potentiel hydrique, conductance stomatique et transpiration.** Les effets du dessèchement du sol sur le flux de sève, le potentiel hydrique des rameaux, la conductance stomatique et du couvert ont été analysés dans 2 placeaux d'un peuplement d'épicéas âgés de 30 ans. L'un des placeaux a été soumis à une sécheresse par couverture du sol, le second ayant été irrigué. Le potentiel hydrique de base des arbres du placeau sec est descendu jusqu'à  $-1,2$  MPa. Dans le placeau irrigué, une faible variabilité de la densité de flux de sève a été observée entre les arbres mesurés, les maxima étant de l'ordre de  $1,2$  à  $1,9 \text{ dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ , ce qui correspondait à environ  $0,5 \text{ mm} \cdot \text{h}^{-1}$ . Dans le placeau desséché, la densité de flux de sève a diminué tout au long de l'été jusqu'à atteindre au minimum  $0,05 \text{ dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$  pour certains arbres, la variabilité entre arbres étant beaucoup plus importante que chez les arbres arrosés. La transpiration ainsi que la conductance stomatique ont fortement diminué avec la sécheresse, la plus grande part de cette réduction ayant été observée lorsque le potentiel hydrique de base est passé de  $-0,4$  à  $-0,6$  MPa. La conductance du couvert, calculée en inversant la formule de Penman-Monteith, a été modélisée au moyen d'un modèle multi-variable prenant en compte les facteurs climatiques et la sécheresse édaphique.

**transpiration / conductance stomatique / conductance de couvert / potentiel hydrique / sécheresse / flux de sève / *Picea abies***

## INTRODUCTION

Norway spruce is one of the most important coniferous forest species used for timber production in Europe. Extensive ecophysiological studies have been done on seedlings and saplings of this species. In contrast, only limited ecophysiological investigations have been reported on adult spruce under field conditions (Schulze *et al*, 1985; Werk *et al*, 1988; Granier and Claustres, 1989; Schulze *et al*, 1989; Cienciala *et al*, 1992), and these studies did not report the long-term effects of limiting soil water conditions.

During the 1980s, a new phenomenon of spruce forest decline occurred in Europe, especially in its western part. Dendrochronological and biogeochemical investigations in the Vosges massif (eastern France) suggested that the decline of spruce in eastern France and western Germany might be mainly related to repeated severe drought events that had occurred since the mid-1970s in these regions (Lévy and Becker, 1987; Probst *et al*, 1990). Further research on spruce decline therefore requires more knowledge of the ecophysiological behaviour of mountain spruce under long-term soil drought.

In a forest ecosystem, transpiration is one of the major water fluxes; its measurement or estimation is of great importance for forest ecologists and hydrologists. In a conifer forest, as demonstrated by Tan *et al* (1978) and Jarvis and McNaughton (1986), transpiration is mainly controlled by vapour pressure deficit (VPD) and stomatal conductance. At the stand level, canopy conductance is considered to be the integration of all the stomatal (including the boundary layer) conductances in the canopy. If transpiration and climatic variables are known over the same time-scale, canopy conductance can be derived from the Penman–Monteith equation (Monteith, 1973). However, with this approach, the key problem is to determine stand transpiration. In this study, we estimate canopy transpiration from the measurement of xylem sap flow with a method suitable for adult forest trees.

In 1990, in the framework of the French Forest Decline Research Program (DEFORPA), extensive ecophysiological investigations were undertaken in a *Picea abies* stand at the Aubure catchment area in the Vosges with the following objectives: 1) to examine forest canopy transpiration and stomatal behaviour under long-term soil

water deficit, as well as the sensitivity of spruce to soil drought (for this point, a comparison of mountain- and plain-growing Norway spruces was carried out); 2) to analyse and model the seasonal variation of canopy conductance under water constraint; and 3) to characterise the alteration of hydraulic conductance on the soil-leaf pathway and monitor the occurrence of xylem cavitation under intensive drought. This paper reports results from the investigation into the first two points; the hydraulic functioning of spruce will be reported in a forthcoming paper.

## METHODS

### Study site

The study site was located on the southern slope of the Aubure catchment area at a mean elevation of 1 050 m. This catchment is situated on the eastern side of the Vosges mountains, France (7°15'E, 48°12'N) and lies on a base-poor gran-

ite bedrock. Annual rainfall is about 1 500 mm and the annual average air temperature is 6°C (Viville *et al*, 1987). A detailed description of the catchment can be found in Probst *et al* (1990). The spruce stand is a dense, 30-year-old plantation, whose main characteristics are presented in table I. Projected leaf area index (LAI) was estimated through 2 independent methods: 1) the relationship between sapwood area and leaf area (Oren *et al*, 1986) gave a value of 5.6; and 2) direct sampling and measurement of needle dry weight (Le Goaster, 1989) gave 6.1.

Two adjacent plots (water stressed (dry) and control (watered)) were selected in autumn 1989. A 12-m-high scaffolding tower was set up in each plot. In the dry plot (30 trees) water was withheld by a surrounding trench (1 m deep) and a plastic roof extending 2 m above soil surface, from July 10 to September 7 1990. Because a natural drought occurred in this region during the experiment, the watered plot was irrigated 6 times (total 58 mm) in July and August 1990.

### Sap flow and stand transpiration

Xylem sap flux density ( $F_d$ ,  $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ ) was measured using 2-cm-long continuously heated

**Table I.** Characteristics of the studied stand and trees.

Treatment	Number of trees	Height (m)	Circumference at 1.3 m (cm)	Sapwood area ( $\text{dm}^2$ )	Crown status
Dry	49	11.4	42.5	0.996	Codominant
	63	13.3	58.0	1.944	Dominant
	64	13.2	46.5	1.240	Dominant
	66	11.8	44.0	1.087	Codominant
Watered	58	13.5	48.0	1.332	Dominant
	59	14.9	71.2	2.752	Dominant
	60	12.0	50.6	1.491	Codominant
	71	12.8	45.5	1.179	Codominant

Age (years): 30; mean height (m): 12.6; mean circumference (cm):  $49.1 \pm 10.1$ ; stem basal area ( $\text{m}^2 \cdot \text{ha}^{-1}$ ): 45.9; sapwood area ( $\text{m}^2 \cdot \text{ha}^{-1}$ ): 31.9; density (stems  $\cdot \text{ha}^{-1}$ ): 2 343; projected leaf area index ( $\text{m}^2 \cdot \text{m}^{-2}$ ): 5.6 (6.1) (2 independent estimates (see text)); altitude (m): 1 050; slope: 15°; slope orientation: SE; mean annual precipitation (mm): 1 500; mean annual air temperature (°C): 6.

sap flowmeters (Granier, 1985, 1987) on 4 trees from each plot, from June to mid-October 1990. The sensors were connected to a datalogger (Campbell Ltd, 21X); measurements were taken every 10 s and hourly means were stored for further processing.

Total sap flow ( $\text{dm}^3 \cdot \text{h}^{-1}$ ) was calculated for each tree by multiplying  $F_d$  by the sapwood cross-sectional area ( $SA$ ,  $\text{dm}^2$ ) of the trees at the sensor level.  $SA$  was estimated using a relationship between tree circumference ( $C$ ) and  $SA$ , established from a sampling of cores on the surrounding trees (Granier, 1985; Lu, 1992):

$$SA (\text{dm}^2) = 0.00047C (\text{cm})^{2.052}$$

$$r^2 = 0.973 \quad [1]$$

Hourly stand transpiration ( $T$ ,  $\text{mm} \cdot \text{h}^{-1}$ ) was computed as:

$$T = SA_T \sum (F_{di} \cdot p_i) \quad [2]$$

where  $SA_T$  was the plot sapwood area per unit of ground area ( $31.9 \text{ m}^2 \cdot \text{ha}^{-1}$ ),  $F_{di}$  the mean sap flux density of trees in the class of circumference  $i$ ,  $p_i = SA_i / SA_T$ , and  $SA_i$  the sapwood area of the trees in the class of circumference  $i$ ; 3 classes were used: dominant trees ( $C \geq 55 \text{ cm}$ ); codominant ( $40 \leq C < 55 \text{ cm}$ ); and intermediate plus suppressed trees ( $C < 40 \text{ cm}$ ).

The characteristics of the studied trees are shown in table I. Daily plot sap flow ( $\text{mm} \cdot \text{d}^{-1}$ ) was calculated as the total of the hourly values.

### Twig water potential

Twig water potential was measured twice a month on 3 one-year-old twigs from each of the studied trees (8 sap flow measured trees plus 2 additional trees from the dry plot), using a pressure chamber. Twigs were sampled in the upper third part of the crown just before dawn (predawn water potential,  $\Psi_{pd}$ ) and at 12:00 solar time during sunny days (midday water potential,  $\Psi_m$ ).

Throughout the study period, 2 trees in each plot (No 66 and 49 from the dry plot; No 59 and 71 from the watered plot) were selected for extensive measurements of diurnal courses of twig water potential. These trees were chosen for the easy access to their crown from the towers.

### Stomatal conductance

Midday stomatal conductance ( $g_s$ ) was measured between 12:00 and 13:00 solar time on 7 sunny days (days 206, 213, 214, 220, 235, 255 and 284) throughout the growing season using a Li-Cor 1600 porometer (Lincoln, USA). Four exposed sun twigs and 4 exposed shade twigs were selected in the upper half of the crown of the 4 extensively measured trees.

### Climatic measurements

Climatic factors above the stand (global radiation, relative humidity, air temperature and wind speed) were measured hourly in a weather station 500 m from the stand. Incident rainfall and throughfall were measured weekly in a cutting and in the watered plot, respectively.

Maximum transpiration ( $TM$ ,  $\text{mm} \cdot \text{h}^{-1}$ ) was calculated hourly from the climatic data using the Penman–Monteith equation:

$$TM = \frac{s(R_n - G) + \rho C_p \text{VPD } g_a}{\lambda [s + \gamma(1 + g_a/g_{cm})]} \quad [3]$$

where:

$s$ : rate of change of saturation vapour pressure ( $\text{Pa} \cdot \text{C}^{-1}$ )

$R_n$ : net radiation above stand ( $\text{W} \cdot \text{m}^{-2}$ )

$G$ : rate of change of heat in the biomass, plus heat in the soil ( $\text{W} \cdot \text{m}^{-2}$ )

$\rho$ : density of dry air ( $\text{kg} \cdot \text{m}^{-3}$ )

$C_p$ : specific heat of dry air at constant pressure ( $\text{J} \cdot \text{kg}^{-1} \cdot \text{C}^{-1}$ )

VPD: vapour pressure deficit (Pa)

$g_a$ : aerodynamic conductance ( $\text{cm} \cdot \text{s}^{-1}$ )

$g_{cm}$ : maximum (non-limiting soil water) canopy conductance ( $\text{cm} \cdot \text{s}^{-1}$ )

$\lambda$ : latent heat of vaporisation of water ( $\text{J} \cdot \text{kg}^{-1}$ )

$\gamma$ : psychrometric constant ( $\text{Pa} \cdot \text{C}^{-1}$ )

In this study, heat flow in the soil was not measured but was assumed to be negligible.  $R_n$  was calculated as 75% of global radiation (unpublished data, from a previous experiment in a spruce stand near Nancy, France). Rate of storage of heat in biomass was calculated from the

above-ground estimated biomass and from hourly changes in air temperature (Stewart, 1988). Aerodynamic conductance ( $g_a$ ) was calculated using the logarithmic equation of Monteith (1973) from wind speed and mean height of the stand (12.6 m). Daily TM ( $\text{mm}\cdot\text{d}^{-1}$ ) was then calculated as the cumulated values of hourly TM.

The maximum canopy conductance ( $g_{cm}$ ) was modelled. It was first calculated hourly from sap flow (in both plots) and climatic data during the beginning of the measurement period (days 164 to 190) under non-limiting soil water conditions, using equation [3]. It was assumed that vapour flux was equal to the stand sap flow scaled up from the trees sap flow, as in Cienciala *et al* (1992). The first tests have shown a 1 h time lag between sap flow and simulated TM. Thus, maximum canopy conductance was recomputed from sap flow measured over hour (h) and climatic factors measured over hour (h - 1). A multiple regression was made on hourly daylight data over the period of days 165 to 190, using a non-linear model close to the equation proposed by Lohammar *et al* (1980):

$$g_{cm} = (R_g / (R_g + 310)) (5.34 - 1.33 \ln(\text{VPD})) \quad [4]$$

$$r^2 = 0.70$$

with  $g_{cm}$  in  $\text{cm}\cdot\text{s}^{-1}$ ,  $R_g$  in  $\text{W}\cdot\text{m}^{-2}$ , and VPD in hPa.

In a forest stand,  $g_{cm}$  can be considered in the first approximation as the average of leaf stomatal conductances over the entire canopy:

$$g_{cm} = g_s \text{LAI} \cdot 2.6 \quad [5]$$

where  $\text{LAI} \cdot 2.6$  is the developed leaf area index of the stand (Oren *et al*, 1986).

### Additional experiment

Another experiment has been undertaken previously near Nancy, France ( $6^\circ 14' \text{E}$ ,  $48^\circ 44' \text{N}$ , elevation 250 m) on a 21-year-old Norway spruce plantation. The stand density was 4 200 stems  $\cdot \text{ha}^{-1}$ , average tree circumference 31.3 cm, and average tree height 11.3 m. The soil was a Gleyic luvisol developed on loam. This experiment was described by Granier and Claustres (1989). Sap flow and xylem water potential measurements were performed on 5 trees from different crown classes, by means of the same technique.

## RESULTS

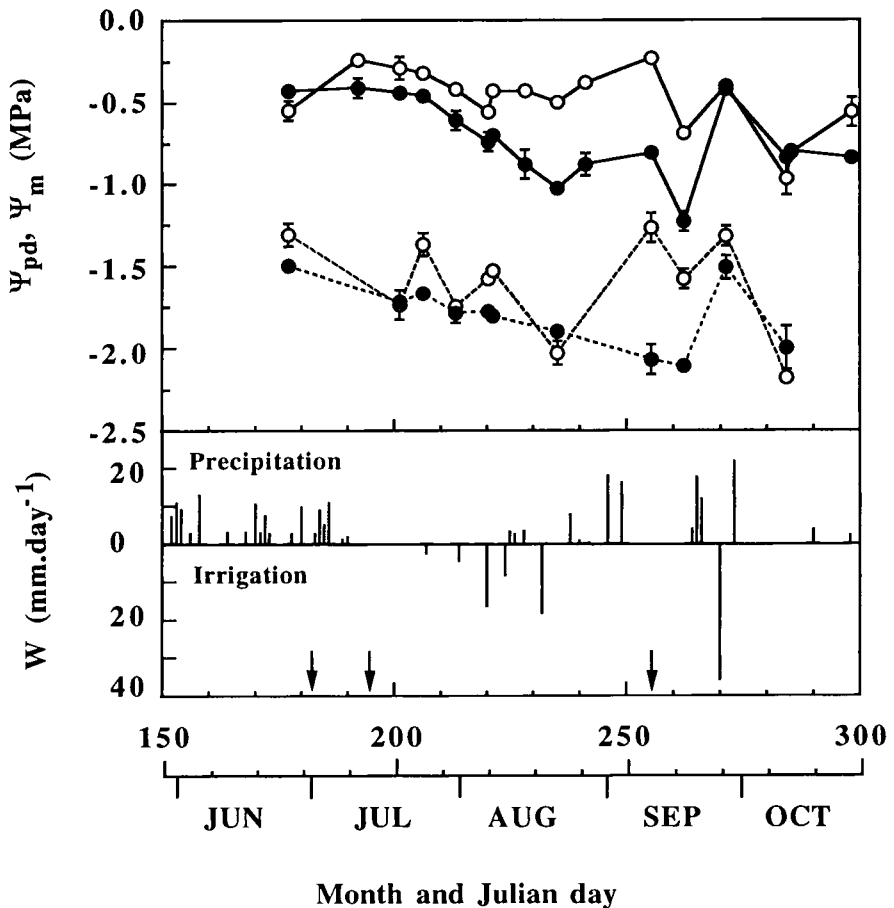
### Twig water potential variations

The year 1990 was characterised by a relatively dry spring followed by an exceptionally dry summer and autumn (Dambrine *et al*, 1992).

Figure 1 shows the seasonal course of average predawn ( $\Psi_{pd}$ ) and midday water potential ( $\Psi_m$ ) of trees in the dry and watered plots. Before the roof was put in place, when the soil was well-watered, the  $\Psi_{pd}$  values in watered and dry plots were  $-0.55$  and  $-0.45$  MPa, respectively, on day 176. Later, a slight difference (about 0.15 to 0.20 MPa) was noticed between both plots, probably due to the trench which immediately provoked a decrease in soil water potential in the dry plot, as was also reported by Biron (1994) from tensiometer measurements. During the following drier and warmer period (days 190 to 238),  $\Psi_{pd}$  and  $\Psi_m$  in both plots first decreased gradually and concurrently until the beginning of the August. Afterwards, due to irrigation in the watered plot (especially on days 220, 225 and 233), the  $\Psi_{pd}$  of the watered plot increased and remained relatively stable around  $-0.4$  MPa. In contrast,  $\Psi_{pd}$  of the dry plot continued to decrease gradually to about  $-1.0$  MPa, and then slightly increased due to several rainfall events from mid-August to mid-September.

After the removal of the roof (September 15),  $\Psi_{pd}$  continued to decrease in both plots in the absence of rainfall and irrigation. At this time, trees in the dry plot were exposed to the most severe drought observed in this experiment ( $\Psi_{pd}$  and  $\Psi_m$  were  $-1.2$  and  $-2.0$  MPa, respectively).

Variations of  $\Psi_m$  progressed in parallel with  $\Psi_{pd}$ , with a difference of about 1.0 MPa. Except for 1 day (day 235), the trees in the dry plot revealed a more negative  $\Psi_m$  than those in the watered plot.

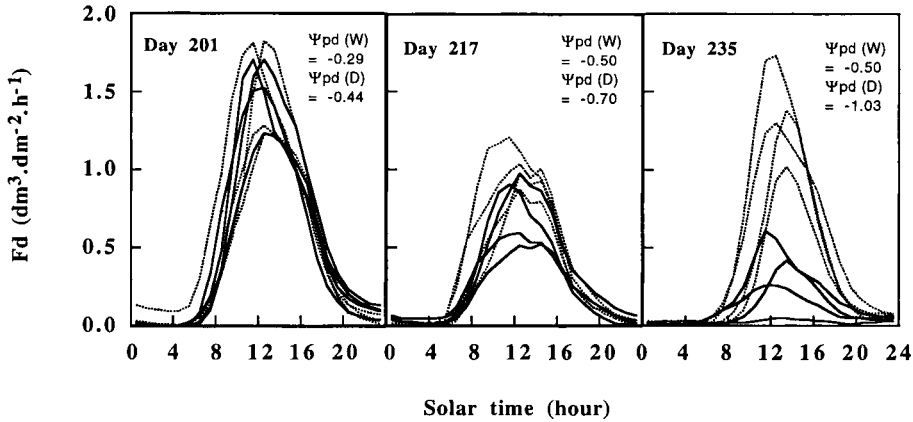


**Fig 1.** Seasonal variations in twig predawn ( $\Psi_{pd}$ , full lines) and minimum water potential ( $\Psi_m$ , dotted lines) in the dry plot (closed symbols) and in the watered plot (open symbols). Rainfall and irrigation in the watered plot (except the last irrigation = 40 mm applied to dry plot) are indicated on the bottom part of the figure. The arrows indicate successively the date of the opening of the trench, the installation of the roof and the removal of the roof. Error bars indicate standard deviation.

### Daily variations of sap flux density ( $F_d$ )

Examples of diurnal course of  $F_d$  during 3 bright days over the season are shown in figure 2. On day 201, under high water availability conditions ( $\Psi_{pd} = -0.29$  MPa in the watered plot, and  $\Psi_{pd} = -0.44$  MPa in the dry plot),  $F_d$  courses were very similar,

and between-tree variability was low. Nevertheless, some differences could be noticed. In the morning, the sharp increase in sap flux densities did not occur at the same time for all the trees, and some of them displayed their maxima earlier than others. Throughout the season, the maximum  $F_d$  varied between 1.2 and 1.9  $\text{dm}^3\cdot\text{dm}^{-2}\cdot\text{h}^{-1}$ , according to the trees.



**Fig 2.** Diurnal patterns of sap flux densities ( $F_d$ ) during 3 bright days on the 4 watered (dotted) and the 4 dry (full lines) measured spruce trees. Mean value of predawn water potential in each plot is also indicated ( $\Psi_{pd}(W)$ : watered plot;  $\Psi_{pd}(D)$ : dry plot).

Increasing the soil water deficit induced a gradual decrease in  $F_d$  and the increase in between-tree variability, as shown on days 217 and 235. Under the driest conditions (eg, on day 235), maximum  $F_d$  (mean  $\Psi_{pd} = -1.03$  MPa) dropped to very low values ( $0.05\text{--}0.5\text{ dm}^3\cdot\text{dm}^{-2}\cdot\text{h}^{-1}$ ), while  $F_d$  in the watered trees remained higher, ranging between 1.0 and  $1.75\text{ dm}^3\cdot\text{dm}^{-2}\cdot\text{h}^{-1}$ . It was also observed that the 2 dominant trees in the dry plot exhibited a much lower  $F_d$  than codominant trees, while no relationship between crown status and  $F_d$  was apparent for the watered trees.

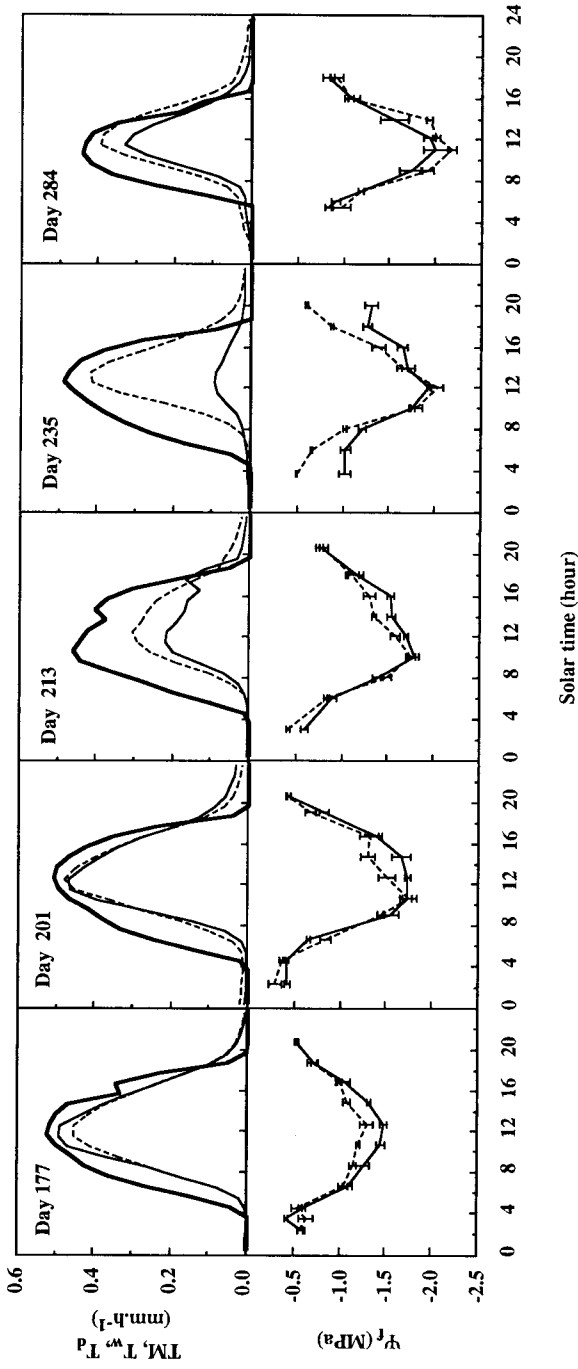
#### **Diurnal and seasonal courses of plot transpiration**

Over the study period, 5 diurnal courses of plot transpiration ( $T_w$ ,  $T_d$ ), maximum transpiration (TM) and average twig water potential ( $\Psi_t$ ) are shown in figure 3, to illustrate the effects of increasing soil drought on plot transpiration. At the beginning of the

season, transpiration values in the 2 plots were similar, with maximal transpiration rates at midday of  $0.43\text{ mm}\cdot\text{h}^{-1}$ . Significant differences between the 2 plots were observed under the higher soil water deficit (days 213 and 235). For example, on day 235, transpiration of the dry plot decreased to less than 25% of that of the watered plot. After irrigation (day 284), transpiration in the dry plot almost recovered to a similar level of the watered plot.

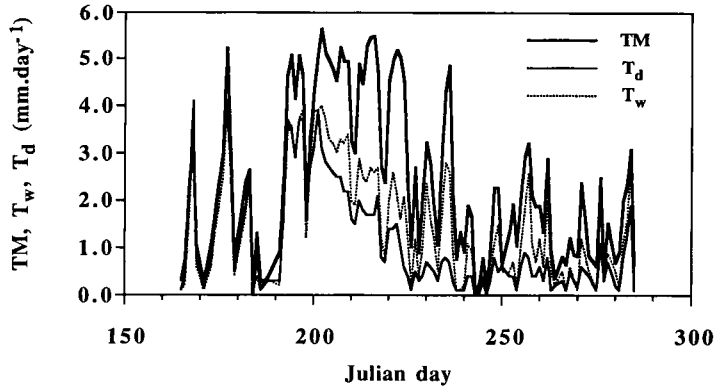
As shown in figure 1, day 235 had one of the lowest  $\Psi_{pd}$ . At this time, comparable values of  $\Psi_m$  (about  $-2.0$  MPa) were observed in the dry and watered plots, suggesting that stomatal closure prevented trees in the irrigated plot from developing more severe water stress. It was also observed that the recovery of twig water potential after sunset was slow under severe water deficits (fig 3, day 235).

Seasonal courses of daily TM,  $T_w$  and  $T_d$  are shown in figure 4. TM was higher during July and August (from days 190 to 235), with maximum values of  $5.5\text{ mm}\cdot\text{d}^{-1}$ ,



**Fig 3.** Top: diurnal patterns of transpiration in watered ( $T_w$ , dotted) and dry ( $T_d$ , full lines) plots and maximum stand transpiration (TM, thick lines). Bottom: diurnal variations of average twig water potential ( $\Psi_t$ ) in the watered (dotted) and dry plot (full lines). Error bars indicate standard deviation.





**Fig 4.** Seasonal variations of dry ( $T_d$ , full line) and watered ( $T_w$ , dotted) plot transpiration as compared to maximum transpiration (TM, thick line).

and during the remainder of the measurement period, it ranged between 1.0 and 4.0  $\text{mm}\cdot\text{d}^{-1}$ . Plot transpiration rates were first at maximum and close to TM from days 160 to 195. After the beginning of July (day 200), plot transpiration decreased in both plots, revealing stomatal closure. Lower transpiration rates were observed in the dry plot where  $T_d$  fell to 0.08  $\text{mm}\cdot\text{d}^{-1}$ . In the watered plot, after an initial decrease, a tendency to stabilise from days 210 to 225 was observed, the maximum transpiration rate being around 2.5  $\text{mm}\cdot\text{d}^{-1}$ .

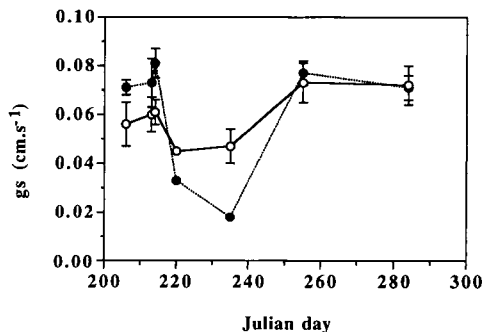
The ratios  $T_d/\text{TM}$  and  $T_w/\text{TM}$  were close to 1 until day 190; afterwards,  $T_d/\text{TM}$  gradually decreased to 0.2 at the end of August, and  $T_w/\text{TM}$  to 0.5, just before irrigation occurred. Over the period from days 165 to 285, the total sums of TM,  $T_w$  and  $T_d$  were 252, 190 and 150 mm, respectively.

### Stomatal control of trees and stand transpiration

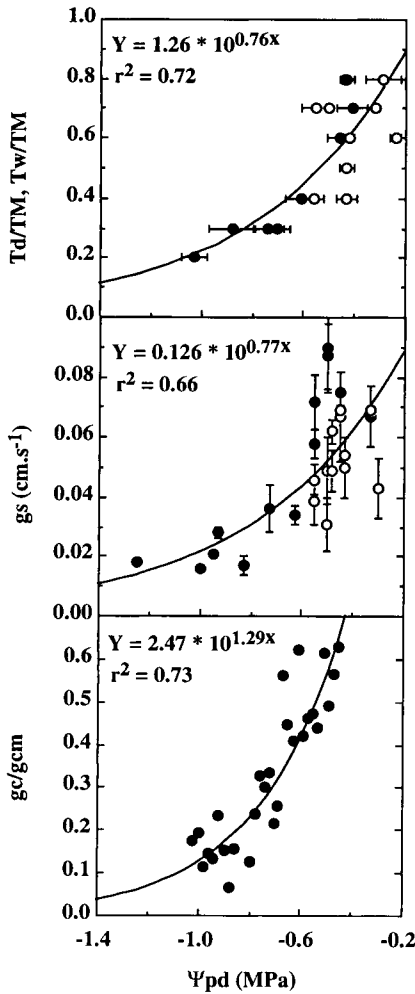
The seasonal course of stomatal conductance ( $g_s$ ) measured at midday is shown in figure 5. Before day 220, stomatal conductances of the watered trees were slightly lower than those in the dry plot, probably resulting from the sampling done at different crown exposures from the towers. A strong decrease of  $g_s$  was observed during July

and August in both plots, but it was more pronounced in the dry plot. Mean  $g_s$  in the dry plot decreased by about 75% from the beginning (0.08  $\text{cm}\cdot\text{s}^{-1}$ ) until mid-August (0.02  $\text{cm}\cdot\text{s}^{-1}$ ), while in the watered plot,  $g_s$  remained quite stable, around 0.05  $\text{cm}\cdot\text{s}^{-1}$ . After the rain at the end of August and the beginning of September, and rehydration of the dry plot, the  $g_s$  in both plots recovered to the pre-stress value.

The decreases in the ratios  $T/\text{TM}$  and  $g_s$  were well correlated with the decrease of predawn water potential in both plots (fig 6). However, most of the decrease in  $g_s$  occurred within a very limited change in predawn water potential (between  $-0.4$  and  $-0.6$  MPa).



**Fig 5.** Seasonal variations of midday stomatal conductance in dry (closed circles) and watered (open circles) trees. Error bars indicate standard deviation.



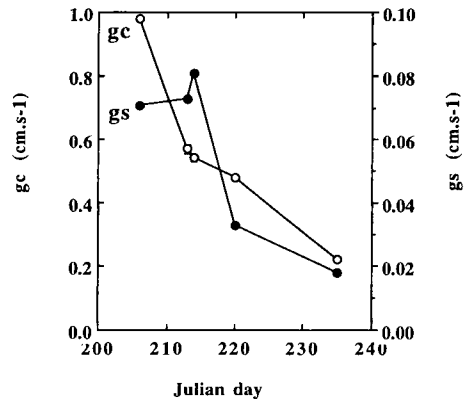
**Fig 6.** Top: ratios of dry ( $T_d$ , closed circles) and watered ( $T_w$ , open circles) plot transpiration to maximum transpiration ( $TM$ ) as a function of predawn water potential ( $\Psi_{pd}$ ). Middle: average midday stomatal conductances in dry (closed circles) and watered (open circles) trees as a function of predawn water potential ( $\Psi_{pd}$ ). The regression curves are fitted with all data from dry and watered trees. Bottom: variation of the ratio of calculated ( $g_c$ ) on modelled maximum ( $g_{cm}$ ) canopy conductance in the dry plot as a function of predawn water potential (period from day 206 to 235, midday values).

At the stand level, drought effects were taken into account in a more general transpiration model than equation [3]. Following Stewart (1988), it was assumed that variations in  $g_c$  could be modelled as the product of a maximum canopy conductance function (under non-limiting soil water conditions, modelled as in equation [4]) and of a function varying between 0 and 1, depending on soil drought. In this study, predawn water potential ( $\Psi_{pd}$ ) was taken as the driving variable. Only midday data were used in order to be compared with stomatal conductance measurements. As previously observed for  $g_s$  variations, figure 6 shows the strong dependence of  $g_c/g_{cm}$  on predawn water potential. A non-linear regression was made between  $g_c/g_{cm}$  and  $\Psi_{pd}$  over the period of dehydration (from day 206 to 235):

$$g_c/g_{cm} = 2.47 \times 10^{1.294\Psi_{pd}}$$

$$r^2 = 0.73$$

Simultaneous variations of  $g_c$  and  $g_s$  (midday values) in the dry plot are shown on figure 7. A good agreement between both courses is observed; the ratio between  $g_c$  and  $g_s$  corresponded approximately to the developed leaf area of the stand, as stated in equation [5].



**Fig 7.** Simultaneous variations in midday canopy conductance ( $g_c$ , open circles) and stomatal conductance ( $g_s$ , closed circles) in the dry plot over the period of increasing soil water deficit.

## DISCUSSION

Under non-limiting water conditions, the maximum hourly sap flux density of the studied trees varied from 1.2 to 1.9  $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$ , which was similar to the values reported in another study on the same species, 1.4–2.2  $\text{dm}^3 \cdot \text{dm}^{-2} \cdot \text{h}^{-1}$  (Granier and Claustres, 1989). Cienciala *et al* (1992) have measured maximum daily sap flux densities of 16  $\text{kg} \cdot \text{dm}^{-2} \cdot \text{d}^{-1}$ , which is in the same range than our values. Between-tree differences in  $F_d$  measured in our study could be attributed to the heterogeneity in crown exposure conditions. We have not found any relationship between  $F_d$  and crown status for the watered trees; dominant trees did not exhibit higher transpiration rates than codominant trees. But under decreasing soil water availability, the  $F_d$  values of the biggest trees were much lower than the  $F_d$  of the codominant trees, indicating a higher soil water depletion by the dominant trees. The minimum  $\Psi_{pd}$  observed in this study was about  $-1.4$  MPa, and  $\Psi_m$  never decreased below  $-2.5$  MPa. This minimum value of  $\Psi_m$  coincided with the threshold of water potential inducing a significant xylem cavitation for this species (Cochard, 1992; Lu, 1992). The mechanism of stomatal closure prevented spruce from xylem dysfunction.

Assessment of the sensitivity of stomata to soil water deficit was one of the principal goals of this study. The relative reduction of  $g_s$  due to the decline of  $\Psi_{pd}$  reported here was comparable to what we observed on spruce growing under similar conditions, in a stand located in central Germany (Lu, unpublished results):  $g_s$  was reduced to about 50% of its initial value when  $\Psi_{pd}$  declined from  $-0.4$  to  $-0.8$  MPa.

Direct comparison of stomata sensitivity to drought between plain and mountain conditions is difficult, because little data are available for spruce growing on the plain. However, comparison between the ratio of stand

transpiration to Penman potential evapotranspiration (T/PET) of the mountain *versus* the plain stands showed a much lower sensitivity to soil drought in the latter than in the former. When  $\Psi_{pd}$  decreased from  $-0.4$  to  $-0.7$  MPa, the reduction of T/PET ratio was only of 20% in the plain stand, compared to 50% in the mountain stand. Nevertheless, we cannot attribute this difference to an intrinsic difference in the stomatal behaviour, because soil and rooting characteristics differ dramatically between both sites. Our mountain stand was located on a shallow sandy soil, with the roots vertically limited by the bedrock. In such a site, soil water depletion develops very rapidly, and therefore a partially desiccated root system could quickly induce stomatal closure, controlled through a biophysical and/or biochemical communication between roots and leaves (Zhang and Davies, 1989; Malone, 1993). Moreover, care must be taken with the use of predawn water potential as a driving variable of stomatal closure. Under field conditions,  $\Psi_{pd}$  does not always seem to be the best indicator of the water stress actually experienced by plants (Reich and Hinckley, 1989; Améglio, 1991). A large decrease of  $g_s$  with only a limited variation in  $\Psi_{pd}$  was observed here, especially for trees in the watered plot (fig 6). This phenomenon has also been reported on the same species by Cienciala *et al* (1994) and in several broad-leaved species such as oak (Bréda *et al*, 1993). When the soil is drying, the upper layers may dehydrate without noticeable change in  $\Psi_{pd}$ . We have shown that during a rainless period, transpiration,  $g_s$  and  $g_c$  are well correlated with  $\Psi_{pd}$  (see fig 6 for the dry plot). However, under variable weather conditions, when some soil layers were dry and others humid (*eg*, after small rain events or irrigation), the implication of  $\Psi_{pd}$  is questionable. So far, there is no clear relationship between  $\Psi_{pd}$  and heterogeneity of water availability in the soil, and it is unclear how the stomatal aperture is controlled in this case. Therefore, more investigations are

needed concerning the significance of  $\Psi_{pd}$  under field conditions.

As demonstrated by McNaughton and Black (1973), for a conifer stand under non-limiting soil water conditions, VPD is the major factor determining tree transpiration, because of a much smaller canopy conductance than aerodynamic conductance, and hence a high degree of coupling between canopies and the atmosphere (Tan *et al*, 1978; Jarvis and McNaughton, 1986; Granier and Claustres, 1989). Except in the morning (when light is limiting), during the course of a day, transpiration is strongly limited by stomatal conductance and its response to VPD variations. Zimmermann *et al* (1988) have indicated the same negative dependence of stomatal conductance to VPD regardless of needle age. Results from the calculation of the canopy conductance (equation [5]) showed that  $g_c$  decreased by about 50% as VPD increased from 0.5 to 1.5 kPa, with  $R_g$  ranging between 500 and 1 000  $W \cdot m^{-2}$ ; in the spruce stand located in the plain, we have observed the same dependence of  $g_c$  to VPD (Granier, unpublished results). As previously emphasised, soil water deficit strongly reduced canopy conductance, decreasing to less than 15% of its initial value as  $\Psi_{pd}$  declined from  $-0.4$  to  $-1.0$  MPa (fig 6).

Maximum midday stomatal conductance values (about  $0.1 \text{ cm} \cdot \text{s}^{-1}$ ) were comparable to data reported in other studies for adult spruce under field conditions (Schulze *et al*, 1985; Claustres, 1987). Canopy conductance variations calculated from sap flow were in good agreement with variations of stomatal conductance (fig 7), even if they were only measured in the upper half of the tree crowns on young needles. Sap flow measured on a representative sample of trees within a stand thus appears to be a valuable method for estimating canopy conductance.

A slow recovery rate of twig water potential after sunset under high water deficit con-

ditions (eg, day 235 in fig 3) was observed. This could be explained by modifications of hydraulic properties within the root zone, where drought induces a high water potential gradient during drought, while water movement is strongly limited by increasing soil hydraulic resistance. Further investigations were done on this question and have shown an important decline of hydraulic conductance, mainly located at the soil-root interface (Lu, 1992).

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